Early Carboniferous (Mississippian) miospore assemblage from Persian Gulf, Southwest Iran

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Abstract

A Mississippian miospore assemblage has been recorded from well A in the Persian Gulf in Iranian offshore, for the first time. Fifteen species of spores have been recognized in the Early Carboniferous assemblage of this well which the most important of them are: *Spelaeotriletes arenaceus, Aratrisporites saharaensis, Spelaeotriletes balteatus, Spelaeotriletes triangulus* and *Radiizonates arcuatus*. Comparing this assemblage with their coevals from the North Africa, Middle East, Western Australia and South America reveals a very close similarity indicating the Gondwanic nature of the assemblage. Also, the Carboniferous strata in this area can be correlated with the Berwath Formation in Saudi Arabia. Based on the palynological data two hiatuses have been recognized at the base and top of the Carboniferous strata.

Keywords: Carboniferous; Mississippian; Gondwana; Miospore; Persian Gulf.

Introduction

The exploratory borehole A, located in the Iranian offshore of Persian Gulf, drilled by the National Iranian Oil Company (NIOC) in 2014 (Fig. 1). Although, the Carboniferous strata are not well known in Zagros and are limited only to a few boreholes in the Southwest of the Zagros fold and thrust belt (Sabouri *et al.*, 2014). They are well known and described from the neighboring Saudi Arabia (Clayton, 1995; Clayton *et al.*, 2000). Therefore, it is important to search and report these rock units due to their paleogeographical position.

The Carboniferous deposits of the Central, Northern and Southwestern Iran have previously been investigated palynologically by Ghavidel– Syooki & Owens (2007), Sabouri *et al.* (2014), Aria–Nasab *et al.* (2016), Sabbaghiyan & Aria– Nasab (2016), Sabbaghiyan (2011; 2016). The first report on the miospore assemblages from the Carboniferous of Southwest Iran (Zagros fold and thrust belt) published by Sabouri *et al.* (2014). They examined cuttings of three boreholes in Fars area and assigned Visean–?early Serpukhovian age to the strata.

Materials and methods

Fourty cutting samples of dark colored, shaly sandstones and shales were collected from well A and prepared in the palynology laboratory of the National Iranian Oil Company (NIOC) for their palynomorph contents. Of these, 16 samples were productive and vielded miospores. Each sample was washed by runnig water to remove drilling mud and 30 grams of each sample were prepared by the method of Traverse (2007). Cold hydrochloric (10%) and hydrofluoric (40%) acids were used to dissolve carbonates and silicates. The residue was neutralized and centrifuged in ZnCl2 (specific gravity 1.9), then sieved with a 15 μ m nylon mesh and mounted on microscopic slides using liquid Canada balsam. Two slides were made from each sample. The microscope slides were examined with an optical microscope and the index spore species were photographed and presented in Plates 1-2. All slides are housed at the Palynology laboratory of the National Iranian Oil Company.

Lithostratigraphy

Carboniferous deposits in well A are 82 m thick and composed of alternations of sandy claystone and sandy shale interbedded with thin bedded sandstone (Fig. 2). These strata here unconformably overlie the Zakeen Formation (Upper Devonian) and are overlain unconformably (an evidence of Hercynian Unconformity) by the Faraghan Formation recently attributed by Spina *et al.* (2018) to the Middle Permian. From lithological point of view, the Carboniferous rock unit in this well can be separated easily from the lower and upper rock units (Fig. 3).

Palynology and palynostratigraphy

In most of the investigated productive slides,

miospores are relatively abundant, diverse and well preserved. Aratrisporites saharaensis, Calamospora liquida, Cingulizonates bialatus, Cyclogranisporites firmus, Densosporites spitsbergensis, Densosporites anulatus, Grandispora maculosa, Indotriradites daemonii, Radiizonates arcuatus, Vallatisporites vallatus, Tricidarisporites serratus, Spelaeotriletes arenaceus, Spelaeotriletes balteatus, Spelaeotriletes triangulus and Spelaeotriletes sp. are the main palynomorphs recorded.



Figure 2. Distribution chart of the selected miospores in well A.



Plate 1. 1–3. *Aratrisporites saharaensis* Loboziak, Clayton & Owens, 1986. 1. Sample 13, 2. Sample 1, 3.sample 16. 4–5. *Spelaeotriletes balteatus* (Playford) Higgs, 1975. Sample 16. 6. *Aratrisporites saharaensis* Loboziak, Clayton & Owens, 1986. Sample 15. 7. *Spelaeotriletes sp.* sample 10. 8–9. *Spelaeotriletes arenaceus* Neves and Owens, 1966. 8. Sample 16, 9. Sample 10. Scale bar represent 25 μm

Most of these species have been reported from other parts of Gondwana such as Saudi Arabia (Clayton, 1995; Clayton *et al.*, 2000), Australia (Playford & Mory, 2017), Niger (Coquel *et al.*, 1995), Algeria (Coquel & Abdesselam–Rouighi, 2000), Libya (Massa *et al.*, 1980; Loboziak & Clayton, 1988), Morocco (Playford *et al.*, 2008) and Brazil (Melo & Loboziak, 2000).

Aratrisporites saharaensis is palaeogeographically a characteristic element of the Early Carboniferous North African "Aratrisporites saharaensis Microflora" (Clayton, 1985; Clayton, 1991) (Fig. 4). This species is known from late Visean deposits of Northern, Eastern and Western Gondwana (Clayton, 1995; Clayton *et al.*, 2000; Coquel & Abdesselam–Rouighi, 2000; Playford & Mory, 2017; Melo & Loboziak, 2000). The species has also been reported from Early Carboniferous (late Tournaisian) deposits of Central Iran (Sabbaghiyan & Aria–Nasab, 2016; Aria–Nasab *et al.*, 2016; Sabbaghiyan, 2016) and Brazil (Playford *et al.*, 2012; Loboziak *et al.*, 1998).

The presence of this index spore in well A can show the relation between Iran and Gondwanan realm during this period of time.

In Northern Gondwana region *Spelaeotriletes arenaceus* has been recognized from late Visean strata (Playford *et al.*, 2008; Coquel & Abdesselam–Rouighi, 2000; Clayton *et al.*, 2000; Coquel *et al.*, 1995).

Densosporites anulatus has been recorded from the Tournaisian–Moscovian age, in Ambo Formation in Peru (Azcuy & di Pasquo, 2005).



Plate 2. 1. *Grandispora maculosa* Playford & Helby, 1968. Sample 1. 2. *Cyclogranisporites firmus* Jones & Truswell, 1992. Sample 15. 3. *Calamospora liquida* Kosanke, 1950. Sample 5. 4. *Indotriradites daemonii* Loboziak et al., 1999. Sample 15. 5–6. *Radiizonates arcuatus* Loboziak et al., 2000. 5. Sample 10, 6. Sample 7. 7. *Cingulizonates bialatus* (Waltz) Smith and Butterworth, 1967. Sample 9. 8. *Tricidarisporites serratus* (Playford) Sullivan & Marshall 1966. Sample 10. 9. *Densosporites anulatus* (Loose) Smith and Butterworth, 1967. Sample 11. 10. *Densosporites spitsbergensis* Playford, 1963. Sample 10. 11. *Vallatisporites vallatus* Hacquebard, 1957. Sample 15. 12. *Spelaeotriletes triangulus* Neves and Owens, 1966. Sample 15. Scale bar represent 25 μm

Spelaeotriletes triangulus is important zonal species of the late Visean to early Serpukhovian. In northeast Libya and the base of the RT biozone is defined on the first appearances of this species and Prolycospora rugulosa (Clayton et al., 2000; Loboziak & Clayton, 1988). Aratrisporites saharaensis, Cingulizonates bialatus, Densosporites spitsbergensis, Indotriradites dolianitii, Radiizonates arcuatus, Spelaeotriletes arenaceus and Spelaeotriletes triangulus have been reported from Visean deposit in Parnaba Basin, Northern Brazil (Melo & Loboziak, 2000).

Miospores as *Aratrisporites saharaensis*, *Cyclogranisporites firmus, Grandispora maculosa* and *Indotriradites daemonii* have been recorded from middle to late Visean *Grandispora maculosa* biozone of Western Australia (Playford & Mory, 2017).

Spelaeotriletes balteatus has been recorded from Tournaisian age within the Spelaeotriletes balteatus–Rugospora polyptycha (BP) and Spelaeotriletes pretiosus–Raistrickia clavata (PC) biozones of Ireland (Brittain & Higgs, 2007) and Belgium (Higgs, 1996).

Several specimens of *Retispora lepidophyta* were recorded as reworked. Also this species has been recorded as reworked miospore in Carboniferous sediment in Saudi Arabia (Clayton, 1995; Clayton *et al.*, 2000; Owens *et al.*, 2000), Northeastern Brazil (Playford *et al.*, 2012) and Australia (Playford & Mory, 2017).

The palynoflora of the Carboniferous strata in well A show a close similarity to the Early Carboniferous palynofloras reported from other parts of Gondwanan regions indicating the Early Carboniferous (late Tournaisian–late Visean) age for this assemblage.

Discussion

The Carboniferous deposits of Saudi Arabia are represented by the Berwath Formation and only identified in subsurface sections (Powers, 1968). The type locality of the Berwath Formation is situated in well ST–8 in northern Saudia Arabia. In South of the Persian Gulf, the Berwath Formation has been recorded in wells 1, 2, 4–6 (Craigie *et al.*, 2016) and Abu Safah–29 well (Clayton, 1995; Clayton *et al.*, 2000) with late Tournaisian–early Serpukhovian age.

Clayton (1995) and Clayton et al. (2000) identified the Carboniferous miospores of three wells in Saudi Arabia. In Arabian side of Persian Gulf, Clayton et al. (2000) reported several such miospore species as **Aratrisporites** saharaensis, Spelaeotriletes triangulus, Radizonates genuinus. prolycospora rugulosa and Spelaeotriletes arenaceus with late Visean age from upper part of Berwath Formation in Abu Safah-29 well. Also, Hooker et al. (2011) reported two biozone (C4-C5) with late Tournaisian-late Visean age from Berwath Formation in Saudi Arabia.



Figure 3. Lithostratigrapic comparison between the Carboniferous deposits of well A in southwest Iran with Berwath Formation of Abu Safah–29 (Gamma ray after Craigie *et al.*, 2016; Lithology redrawn from Clayton *et al.*, 2000) in SaudiArabia



Figure 4. Distribution of Early Carboniferous (lateVisean) Microflora (after Lopes et al., 2016).

Miospore assemblage of well A can be correlated to these two Biozones. For the first time in this study, Carboniferous deposits were identified in Iranian side of Persian Gulf. The recognized Early Carboniferous assemblage of this well dominantly contains *Spelaeotriletes arenaceus* and *Aratrisporites saharaensis*. Also the index species of Visean such as *Cingulizonates bialatus* and *Grandispora maculosa* are present in this assemblage.

However, all of these miospores more or less are quite similar to those of the Gondwanan realm. Also, the examined strata in well A are similar in fossil content and lithology to Berwath Formation of Saudi Arabia. The Carboniferous deposits in Iranian side of Persian Gulf (well A) are compared with their coeval deposits (Fig. 3) mainly from Arabian side of Persian Gulf (Abu Safah–29). The palynological evidence indicates the disconformity between Zakeen and Carboniferous deposits in this well can be assigned to the period of erosion because the latest Devonian index miospores (e.g. *Retispora lepidophyta*) appear as reworked specimen in carboniferous strata.

Conclusions

Moderate diversity and well-preserved palynofloras have been recorded from Early Carboniferous strata in well A of the Persian Gulf Southwest Iran that dominated by miospores. Based on the stratigraphic value of the recognized miospore taxa and comparison with other parts of the Gondwana a late Tournaisian-late Visean age attributed to the assemblage. Biostratigraphical data indicate two hiatuses in the base and top of the Carboniferous strata. The late Tournaisian-late Visean assemblage from well A in Southwest Iran is very similar to coeval Gondwanan assemblage (Algeria, Saudi Arabia, Libva, Morocco, north Brazil, Australia, Central and North Iran). Also, the examined Carboniferous strata in well A are lithologically similar to the Berwath Formation of Saudi Arabia and they are differentiated from the overlying (Faraghan) and underlying (Zakeen) deposits.

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